

BASIC FLUVIAL GEOMORPHOLOGY

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BASIC FLUVIAL GEOMORPHOLOGY

- Interest of the FluvialGeomorphology in the river restoration projects
- Morphological analysis of the rivers
- The sediments in the fluvial processes
- The river in equilibrium: natural dynamics
- Dominant discharge concept
- Response of rivers to natural/human disturbances

Why the Fluvial Geomorphology is important in river restoration

Use and abuse of rivers

- Physical degradation is one of the most important problems of river ecosystems, affecting their ecological status by means of:
 - Lost of natural forms and processes
 - Unwanted erosion and sedimentation processes
 - Decrease of habitat quality and biodiversity
 - Lost of environmental values of rivers

Why the Fluvial Geomorphology is important in river restoration

- The geomorphological analysis of rivers helps to:
 - Identify problems related to the physical degradation
 - Interpretate causes and consequences
 - Propose alternatives for enhancing and restoring rivers

Examples of physical degradation in rivers

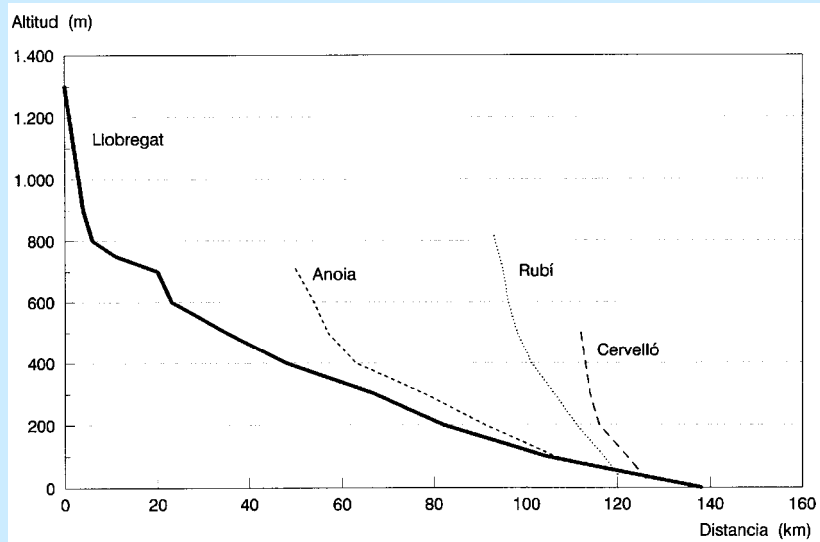


Analysis of Fluvial Forms

FORMS:

- Longitudinal Profile
- Pattern and Sinuosity
- Hydraulic geometry
- Sediments and Bed forms

LONGITUDINAL PROFILE



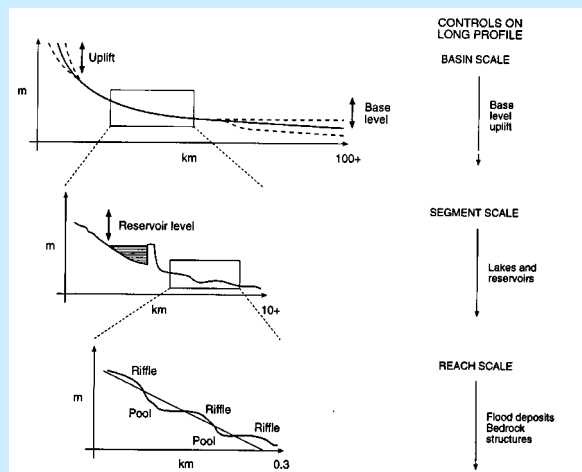
LONGITUDINAL PROFILE IN RIVERS

The longitudinal profile shows the slope of the river, which can be considered at different spatial scales

Watershed (Valley):
Geology and Relief

Fluvial segment:
Dams and local controls

Fluvial reach:
Sediments and bed forms



VARIABLES RELATED TO THE LONGITUDINAL PROFILE OF RIVERS

- The slope of the channel is one of the most important hydraulic variables, determining the hydraulic power and channel stability
- The slope is related to the water velocity and the shear stress
- It is related to the sediment size:

$$S = 18 \left(\frac{d}{A_d} \right)^{0.6}$$

S: Slope (m/m)

A_d : Drainage area (milla²)

d: Medium diameter of sediments (mm)

PATTERN OF RIVERS

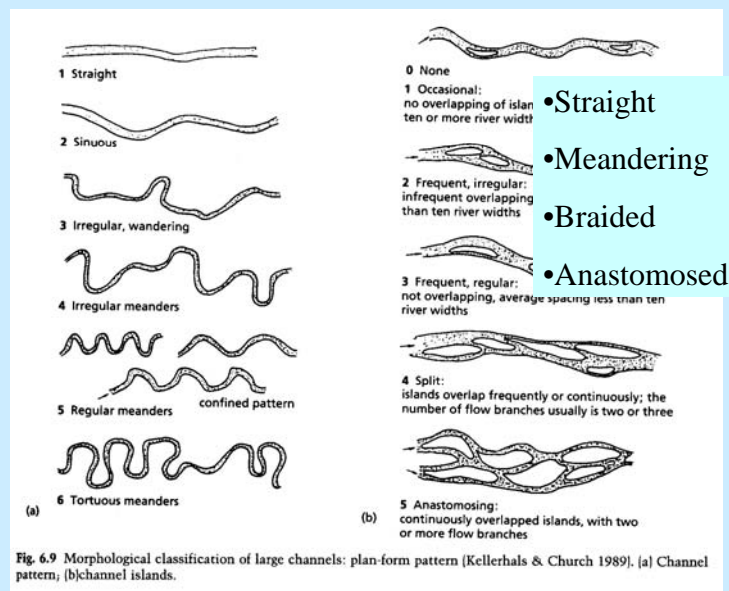


Fig. 6.9 Morphological classification of large channels: plan-form pattern (Kellerhals & Church 1989). (a) Channel pattern; (b) channel islands.

VARIABLES RELATED TO THE PATTERN OF RIVERS

- Slope
- Discharge

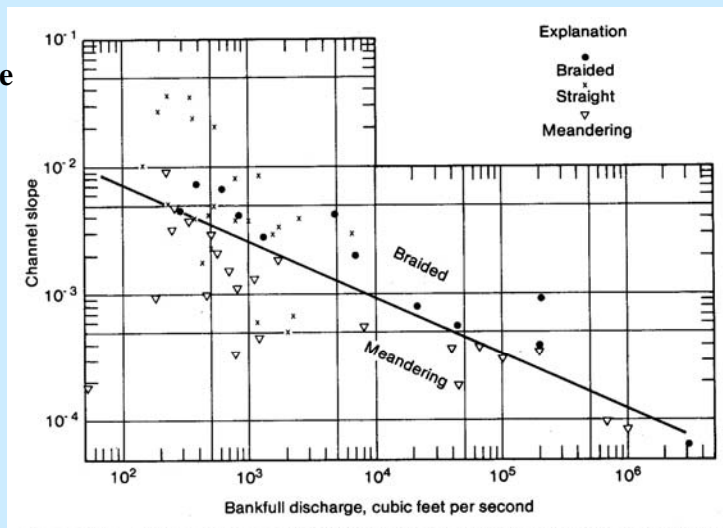
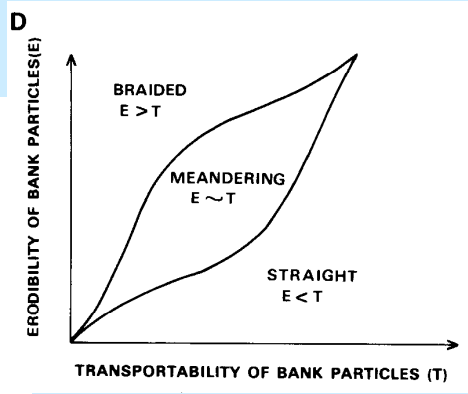
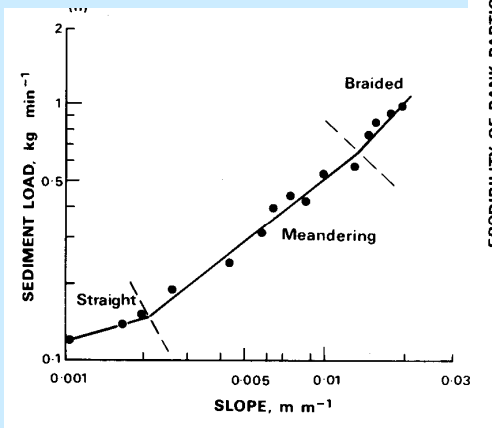


Figure 2.11 Values of slope and bankfull discharge for natural channels as well as a threshold distinguishing braided from meandering rivers (Leopold and Wolman, 1957).

PATTERN OF RIVERS

- Slope
- Sediment load



Bank materials

EXAMPLES OF RIVER PATTERNS

Straight reaches:

LATERAL CONTROLS



EXAMPLES OF RIVER PATTERNS

Meandering rivers

STABLE PATTERN

LOW SLOPE



EXAMPLES OF RIVER PATTERNS

Braided rivers

UNSTABLE PATTERN
SEDIMENT LOAD, HIGH SLOPE



RELATIONSHIP BETWEEN SLOPE, SEDIMENT SIZE AND PATTERN



Valle glacial: Low slope, fine sediments, high sinuosity



Valle fluvial: High slope, coarse sediment, straight pattern (sinuous valley)

CROSS SECTIONS - HIDRAULIC GEOMETRY

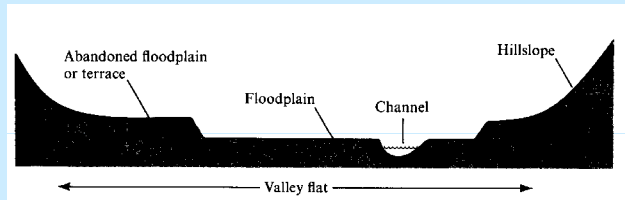
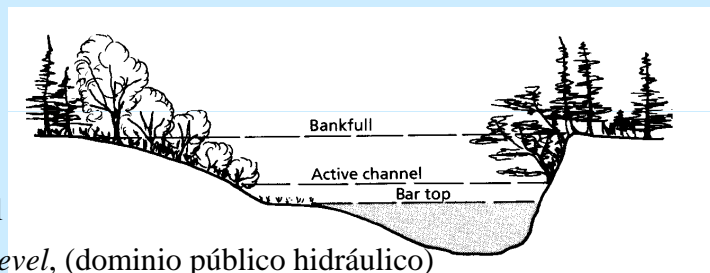


Figure 16-9 Diagrammatic cross section of a valley showing relation of present channel to the floodplain and to a terrace (abandoned floodplain).



- Base level
- Bankfull level, (dominio público hidráulico)
- Floodplain (riberas, llanura de inundación)

HIDRAULIC GEOMETRY

$$\begin{aligned}
 w &= C_a Q^a & Q &= wdu \\
 d &= C_b Q^b & C_a C_b C_c &= 1 \\
 u &= C_c Q^c & a + b + c &= 1
 \end{aligned}$$

At-a-station:

$$\begin{aligned}
 a &= 0,26 \\
 b &= 0,40 \\
 c &= 0,34
 \end{aligned}$$

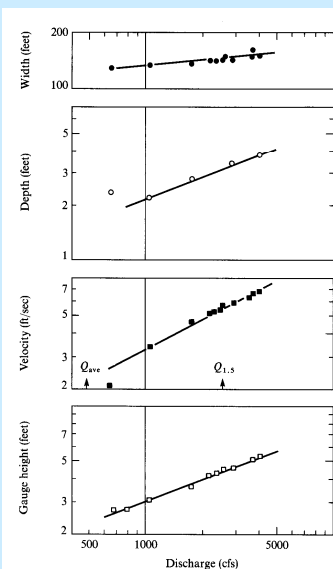


Figure 16-32 At-a-station curves and rating curve for Green River at Warren Bridge, Wyoming.

HIDRAULIC GEOMETRY

$$\begin{aligned}
 w &= C_a Q^a & Q &= wdu \\
 d &= C_b Q^b & C_a C_b C_c &= 1 \\
 u &= C_c Q^c & a + b + c &= 1
 \end{aligned}$$

Downstream, at bankfull:

$$\begin{aligned}
 a &= 0,39 - 0,60 \\
 b &= 0,29 - 0,40 \\
 c &= 0,09 - 0,28
 \end{aligned}$$

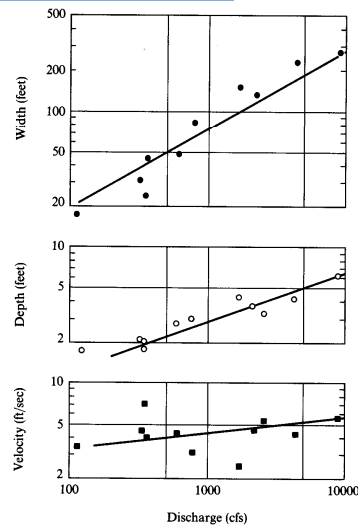


Figure 16-34 Downstream change of width, depth, and velocity with bankfull discharge, upper Green River basin, Wyoming.

SEDIMENTS IN RIVERS

They come from watershed erosion or channel erosion and determine the turbidity of the water, the size of the substratum and the slope of the channel banks

	Transport	Origin
Wash load	Suspension	Watershed erosion Channel erosion (fine particles)
Bed load	Siltation	Channel erosion (coarse particles) Bed erosion

- Cohesive materials
 - Non-cohesive materials
- - Sandy Rivers
 - Gravel-bed Rivers

SANDY RIVERS

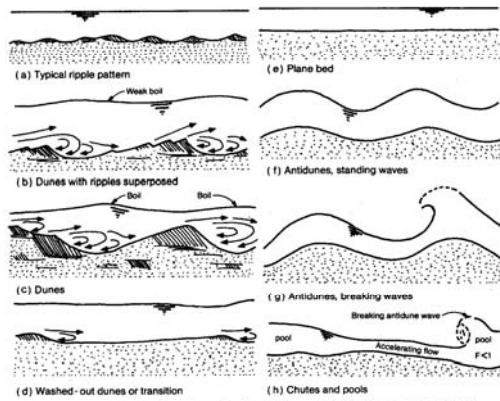


Figure 6.2 Idealized bed forms in alluvial channels (after Simons et al., 1966).

SANDY RIVERS

Bed forms determine the roughness of the channel (Manning's n), which varies according to the discharge

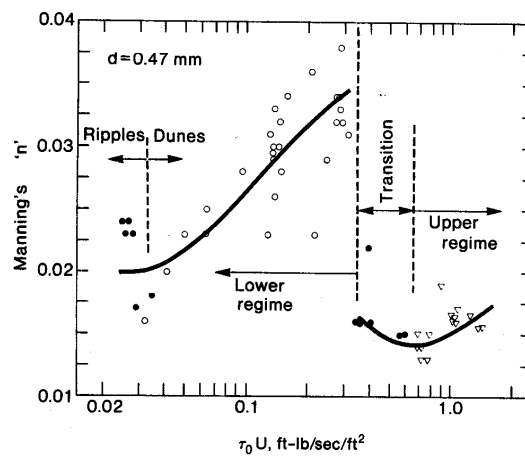
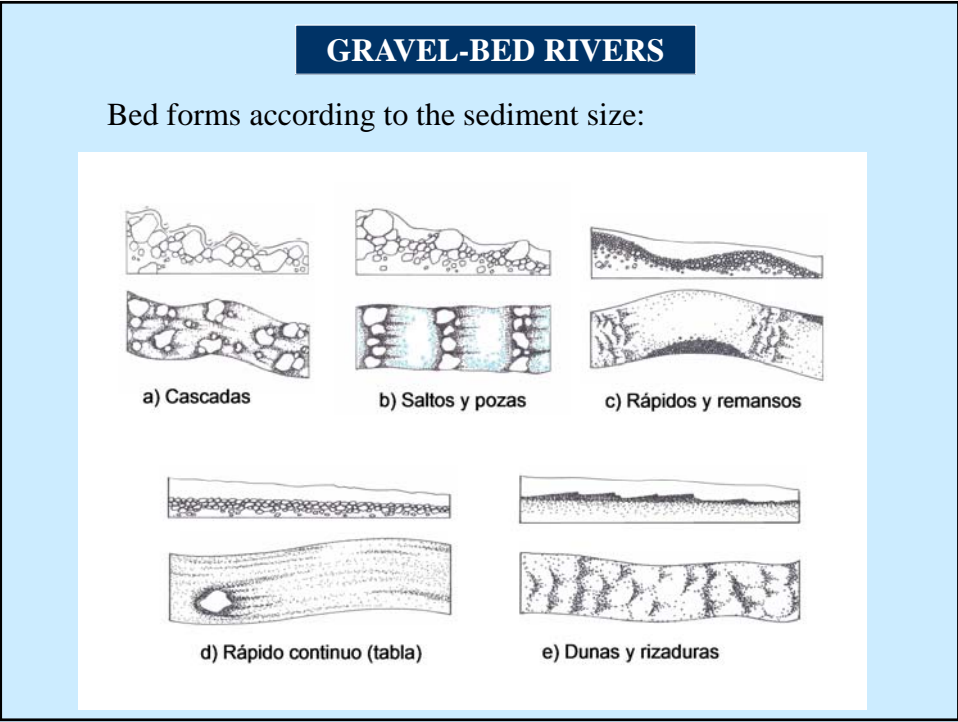
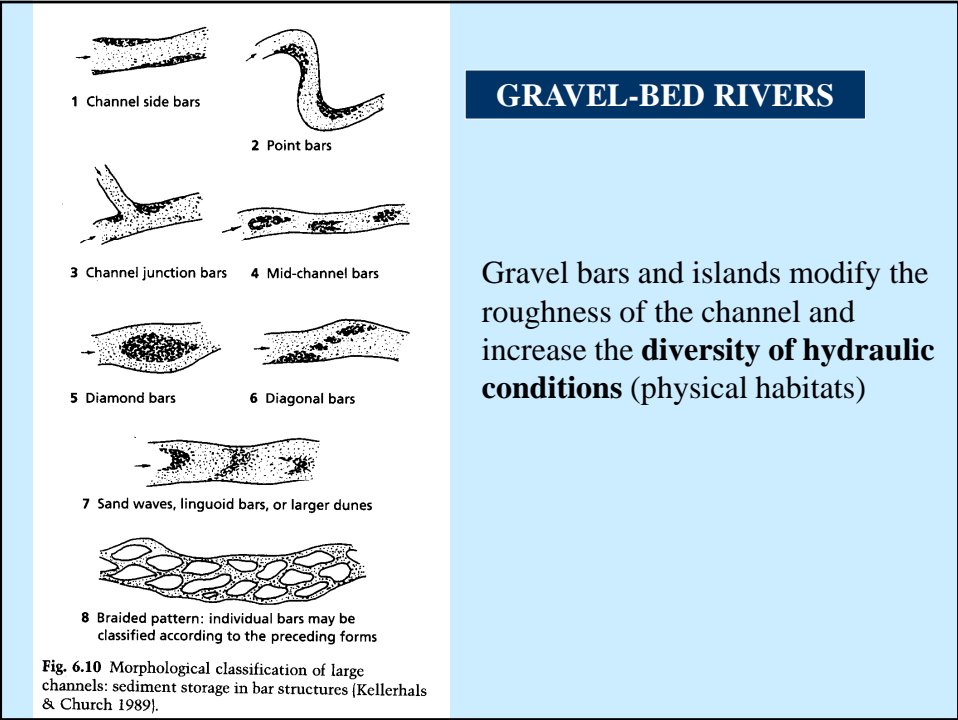


Figure 6.3 Relationship between bed roughness and bed form based on flume data by Guy et al. (1966).



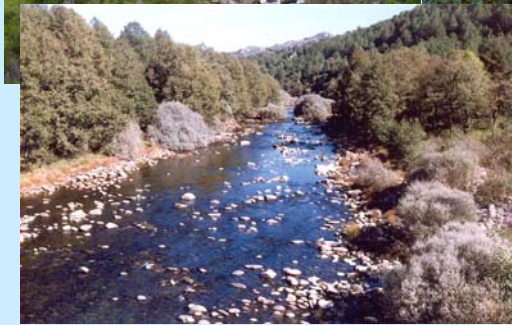
GRAVEL-BED RIVERS

Bed forms according to the sediment size:



Cascade

Steps



Plane bed

GRAVEL-BED RIVERS

Bed forms according to the sediment size:

Pools and Riffles



POOLS AND RIFFLES

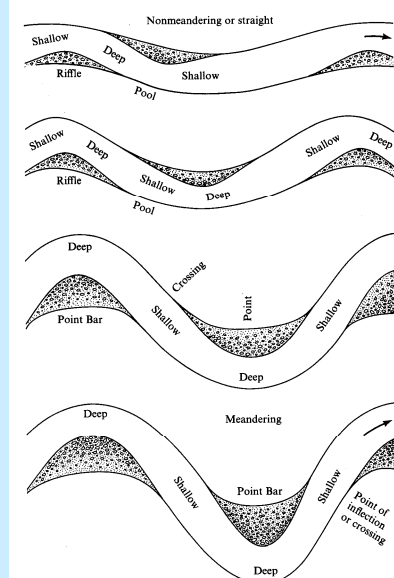
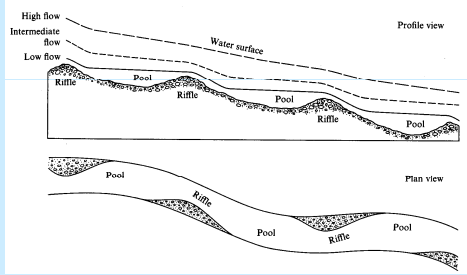
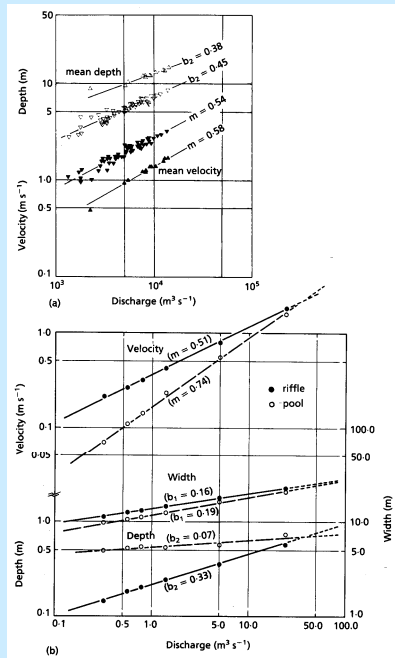


Figure 16-23 Locations of shallow and deep zones in channels of different sinuosity. Riffle bars on alternate banks characterize straight channels, but point bars on convex banks characterize meander bends.

POOLS AND RIFFLES

Fig. 6.13 At-a-station hydraulic geometry. (a) Fraser River, British Columbia at Agassiz (gravel bed ($\nabla \blacktriangledown$)) and at Mission (sand bed ($\triangle \blacktriangle$)); $m = 1 - b_1 - b_2$. (b) Hangover Creek, Queen Charlotte Islands, British Columbia, a boulder-gravel stream of intermediate size, averaged over major morphological units (Hogan & Church 1989)



TYPES OF EQUILIBRIUM

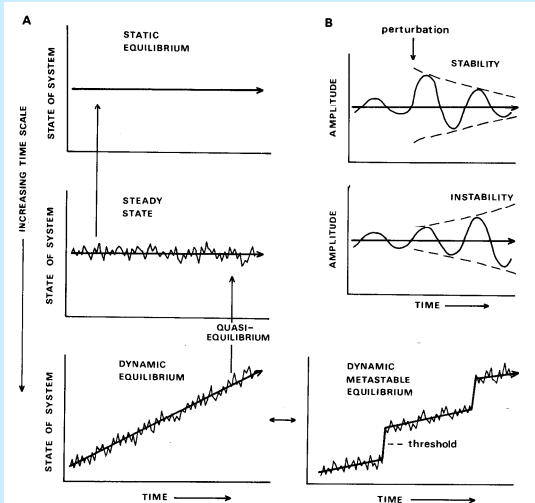
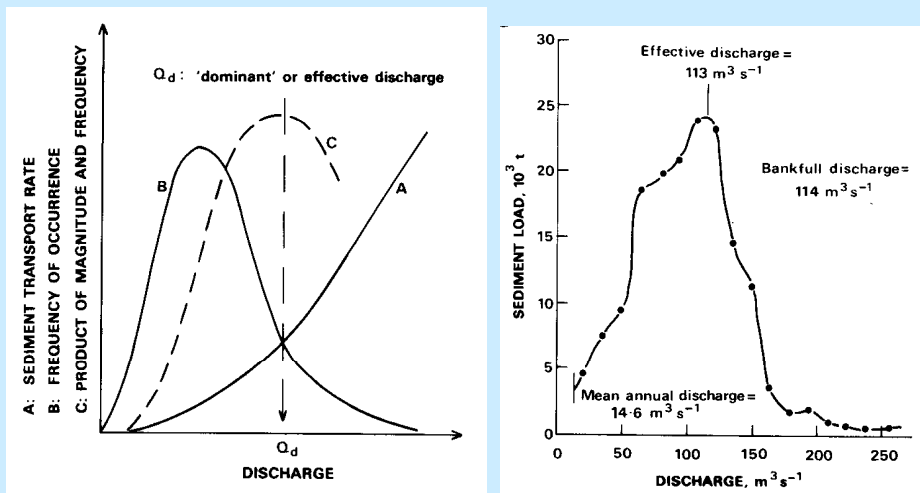


Fig 4.2 A. Diagrammatic representation of types of equilibrium.
 B. Definition sketches of stability and instability in an oscillating mechanical system subject to a perturbation.

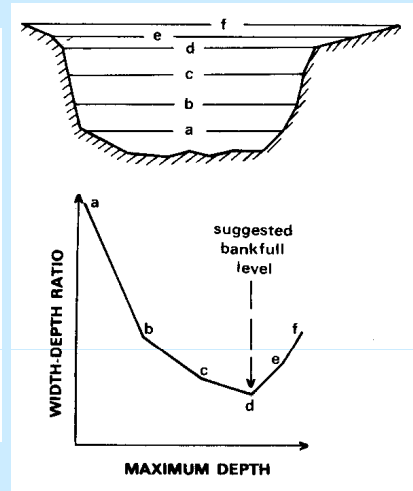
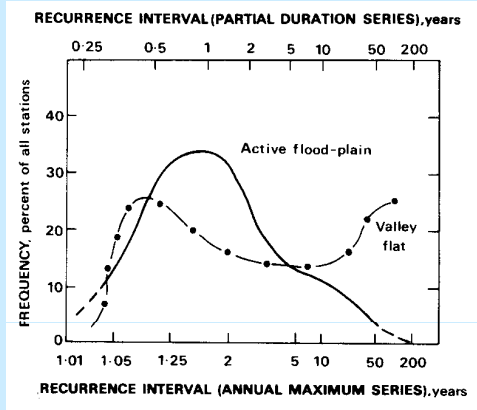
More probable state:

- Uniform distribution of power
- Minimum work (minimum energy lost per unit length)
- Minimum hydraulic power $v \cdot S$

DOMINANT DISCHARGE



DOMINANT DISCHARGE



FLUVIAL RESPONSE

Lane's equation (1955) for predicting fluvial responses:

$$Q * S \approx Q_s * D_{50}$$

FLUVIAL RESPONSE

$$Q * S \approx Q_s * D_{50}$$

Table 1.2 Channel response to natural or man induced changes.
(+ = increase; - = decrease; above line = control; below line = response)

NATURAL Meander cut-offs	$S^+Q_+ \leftrightarrow Q_sD_+$	up	
	$S_+Q \leftrightarrow Q_s^+D_+$	down	
Landslide: Large sustained supply	$S^+Q_+ \leftrightarrow Q_sD_+$	up	
	$S_+Q^- \leftrightarrow Q_s^{++}D_+$	down	
Small limited supply	$S_+Q \leftrightarrow Q_s^+D_+$		

FLUVIAL RESPONSE

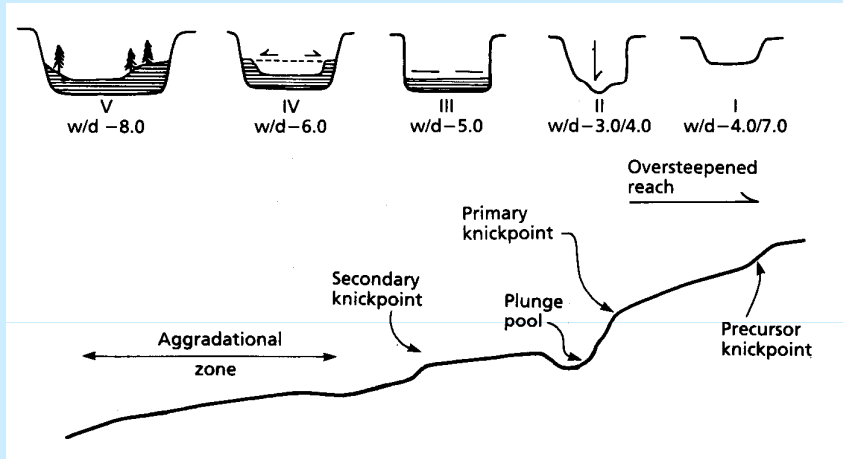
$$Q * S \approx Q_s * D_{50}$$

Table 1.2 Channel response to natural or man induced changes.
(+ = increase; - = decrease; above line = control; below line = response)

ARTIFICIAL Dam construction	$S^-Q^- \leftrightarrow Q_s^-D_+$		
Weir construction	$S^-Q^- \leftrightarrow Q_sD_+$	up	
	$S_+Q_+ \leftrightarrow Q_s^+D_+$	down	
Weir failure	$S^+Q_+ \leftrightarrow Q_sD_+$	up	
	$S_+Q_+ \leftrightarrow Q_s^+D_+$	down	
Channel straightening	$S^+Q_+ \leftrightarrow Q_sD_+$	up	
	$S_+Q_+ \leftrightarrow Q_s^+D_+$	down	
Channel dredging and/or gravel mining	$S^+Q_+ \leftrightarrow Q_sD_+$	up	
	$S^-Q^- \leftrightarrow Q_sD_+$	centre	
	$S_+Q_+ \leftrightarrow Q_s^+D_+$	down	
Interbasin transfers (flows above sediment transport threshold)	$S^+Q_+ \leftrightarrow Q_sD_+$	up	
	$S_+Q^{++} \leftrightarrow Q_sD_+$	down	

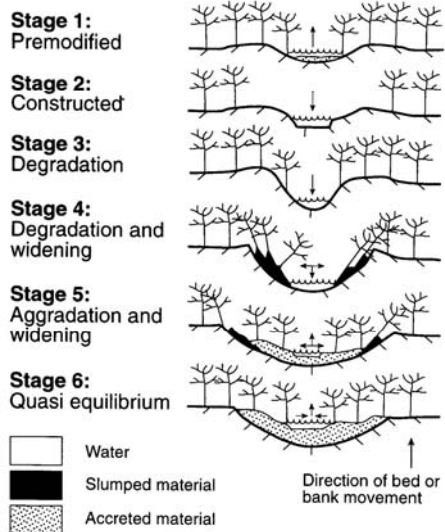
FLUVIAL RESPONSE

INCISED CHANNELS

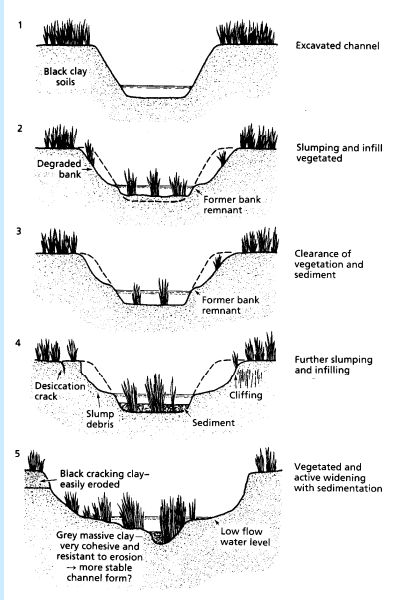


FLUVIAL RESPONSE

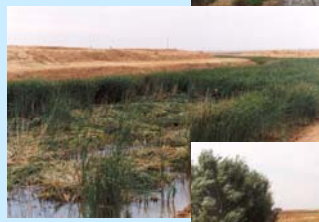
INCISED CHANNELS



FLUVIAL RESPONSE

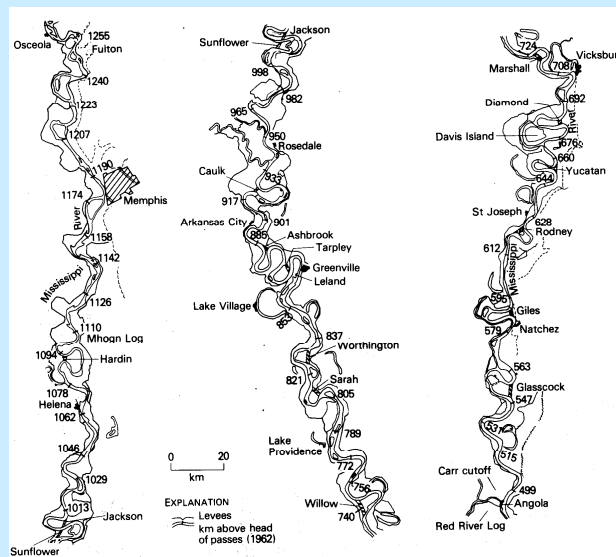


Río Valderaduey



FLUVIAL RESPONSE

Meandering cutoffs in the Mississippi River

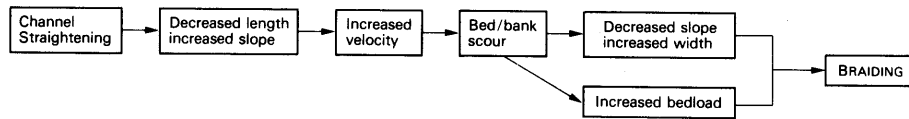


FLUVIAL RESPONSE

Table 11.10 Artificial cutoffs on the Mississippi River, 1932-42. After Winkley (1973)

<i>Cutoff</i>	<i>Year</i>	<i>Shortening (km)</i>
Hardin	1942	27.2
Jackson	1941	14.0
Sunflower	1942	16.7
Caulk	1937	24.5
Ashbrook	1935	18.4
Tarpley	1935	13.8
Leland	1933	15.8
Worthington	1933	6.9
Sarah	1936	8.5
Willow	1934	12.4
Marshall	1934	6.8
Diamond	1933	19.3
Yucatan	1929	15.4
Rodney	1936	9.3
Giles	1933	17.9
Glasscock	1933	17.4
Total		244.3

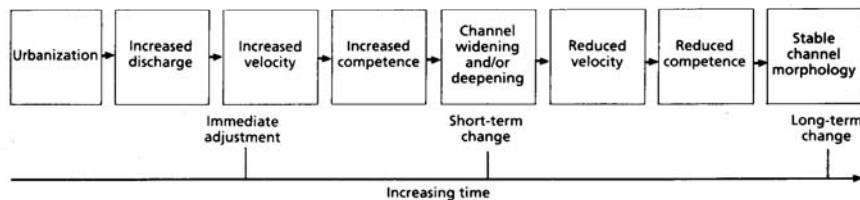
Fig. 11.21 Reaction of a river to a cutoff or channel straightening and its readjustment over time



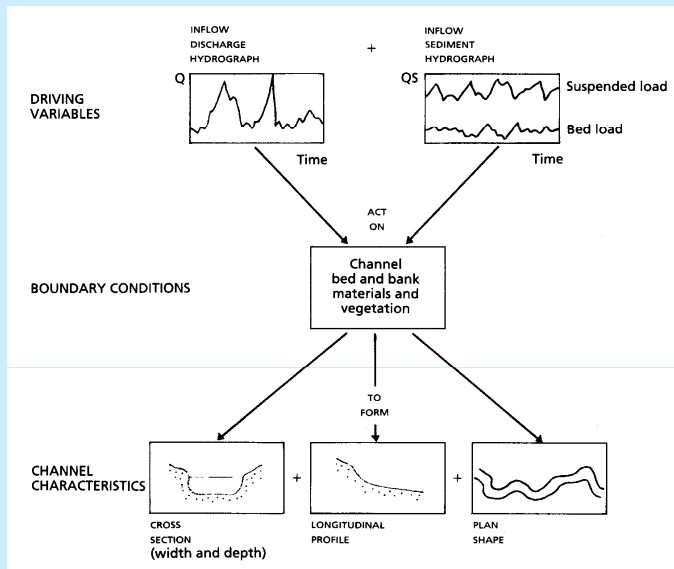
FLUVIAL RESPONSE

Table 11.4 Ratio of discharge after urbanization to discharge before urbanization

<i>Area</i>	<i>Inc. Q_0</i>	<i>Source</i>
Chicago	2.5	Ramey (1959)
Michigan	3.0	Wiitala (1961)
Washington, D.C.	1.8	Carter (1961)
Texas	2-5.0	Van Sickle (1962)
Long Island, New York	2.5	Seaburn (1964)
Mississippi	2-3.5	Wilson (1967)
East Coast	2-3.0	Anderson (1968)
Texas	2-2.5	Espey and Wislow (1974)



FLUVIAL DYNAMICS



•Watershed Hydrology

•Land uses

Human disturbance

•Channel forms

•Fluvial Processes

River Response

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